

Evaporative demand and drought @ PSD: opportunities in early warning, and monitoring.

Mike Hobbins

mike.hobbins@noaa.gov

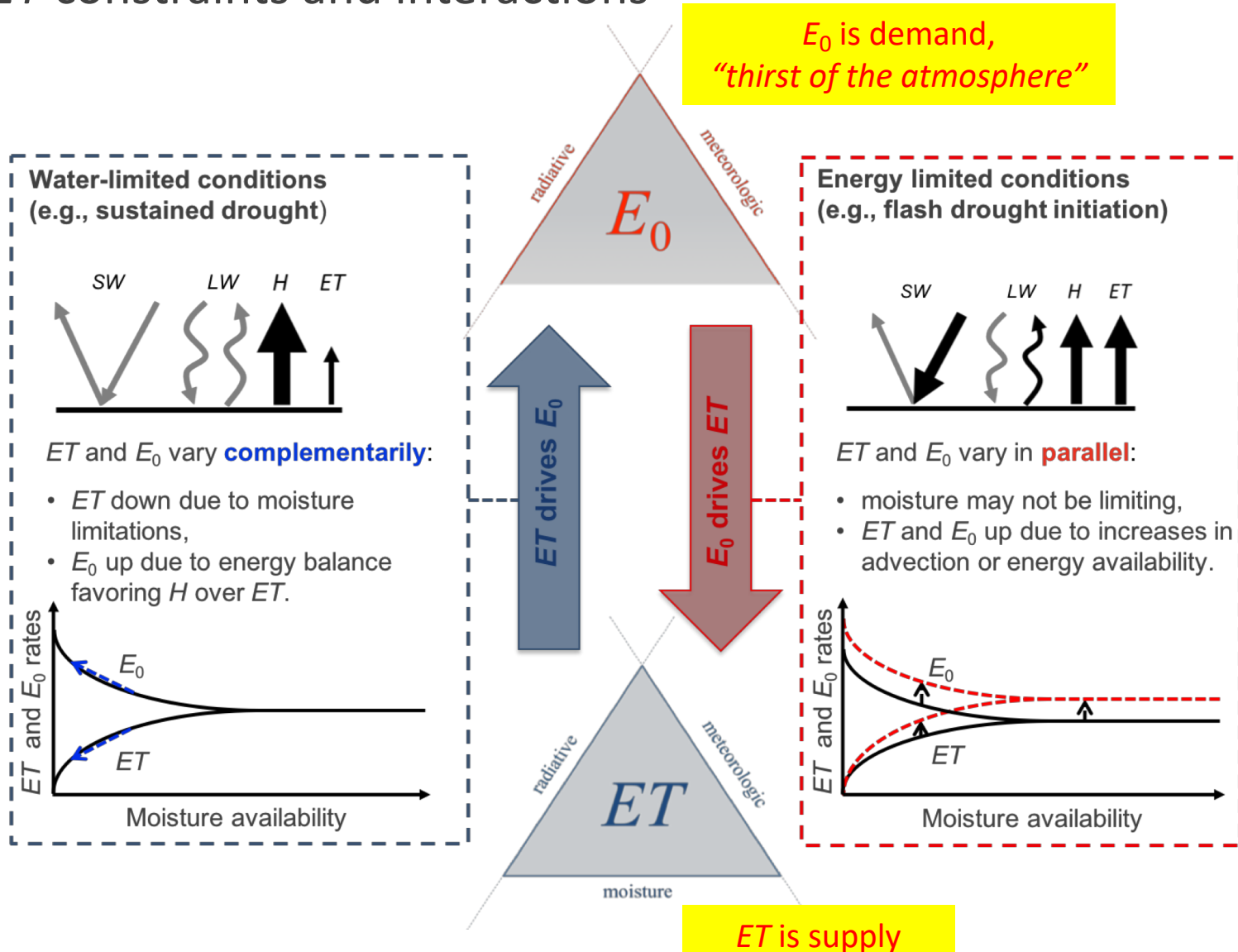
303-497-3092

NOAA-Physical Sciences Division
University of Colorado-CIRES

CREST / NOAA Research and Education Collaboration meeting, August 13-15, 2018



E_0 / ET constraints and interactions



E_0 ➤ much easier to estimate than ET
➤ increases in all drought types.

A demand-side treatment of drought

T = air temperature
 q = specific humidity
 U_2 = wind speed
 R_d = solar radiation

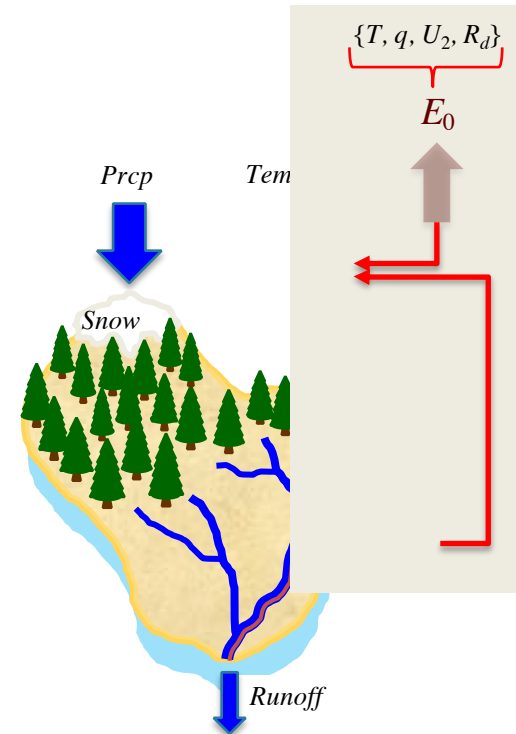
Drought = imbalance of supply to,
and demand for,
surface moisture

Water balance at land surface:

$$\sim f(\text{Prpc}, \text{ET})$$

where ET is more physically driven by:

- surface moisture status,
- evaporative demand (E_0),
 - e.g., Penman-Monteith.



Estimating E_0 from reference ET

Reference ET from FAO-56
(Penman-Monteith):

$$ET_0 = \underbrace{\frac{0.408\Delta}{\Delta + \gamma(1 + C_d U_2)} (R_n - G) \frac{86400}{10^6}}_{\text{Radiative forcing (sunshine, } T)} + \underbrace{\frac{\gamma \frac{C_n}{T}}{\Delta + \gamma(1 + C_d U_2)} U_2 \frac{(e_{sat} - e_a)}{10^3}}_{\text{Advection forcing (wind, humidity, } T)}$$

Reference crop specified:

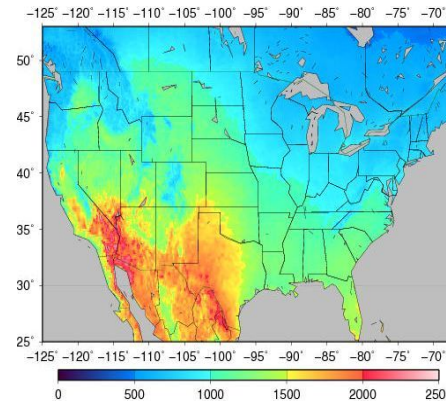
- 0.12-m grass or 0.50-m alfalfa
- well-watered, actively growing,
- completely shading the ground,
- albedo of 0.23.

Drivers:

- T , temperature at 2 m
- q , specific humidity at surface
- R_d , downward SW at surface
- U_2 , wind speed at 2 m

Two reference ET reanalyses

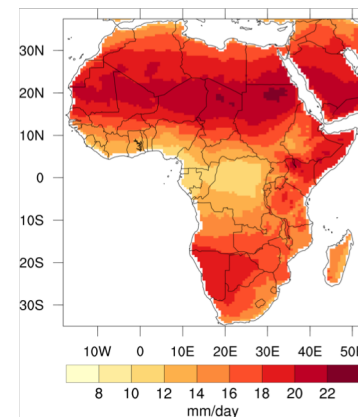
CONUS



Drivers from **NLDAS-2**:

- daily, Jan 1, 1979 – present
- latency ~ 5 days
- **0.125° lat x lon,**
- **CONUS+ (25°N to 53°N)**

GLOBAL



Drivers from **MERRA2**:

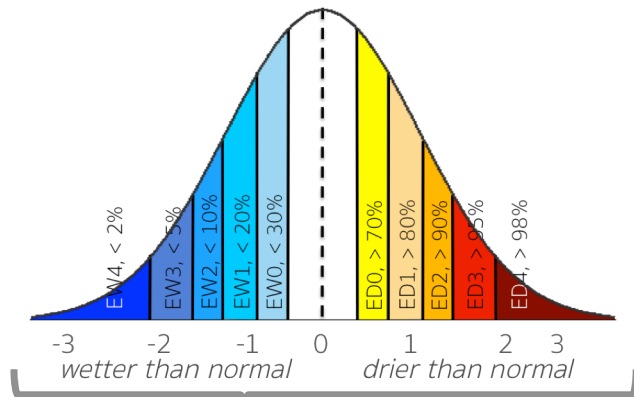
- daily, Jan 1, 1980 – present
- latency ~ 10 days
- **0.5° lat x 0.625° lon**
- **global**

Mean daily E_0 (mm), 1981-2010

Evaporative Demand Drought Index (EDDI)

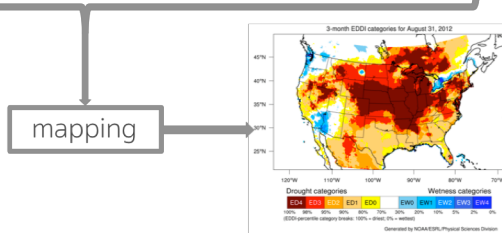
Hobbins et al., *JHM*, 2016

McEvoy et al., *JHM*, 2016

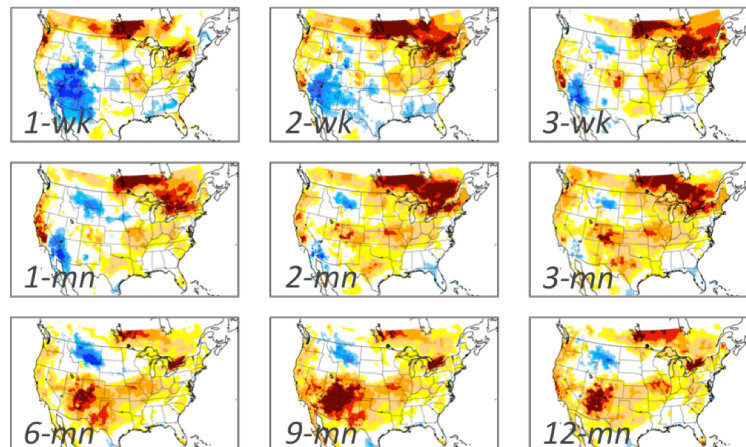


<https://www.esrl.noaa.gov/psd/eddi/>

or search for “EDDI NOAA”

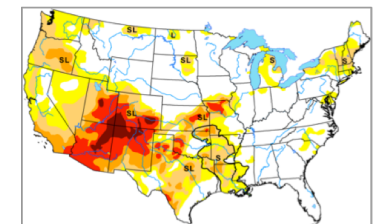


EDDI,
July 17, 2018



EDDI is multi-scalar:
optimal timescale(s) specific to
region, season, and sector.

US Drought Monitor,
July 17, 2018



Intensity:
D0 Abnormally Dry
D1 Moderate Drought
D2 Severe Drought
D3 Extreme Drought
D4 Exceptional Drought

EDDI – cross-sectoral monitoring



AGRICULTURAL
DROUGHT

- soil moisture
- grazing health
- ET

HYDROLOGIC
DROUGHT

- streamflow
- snowfall



FIRE-RISK
MONITORING

- weather
- fuel loads

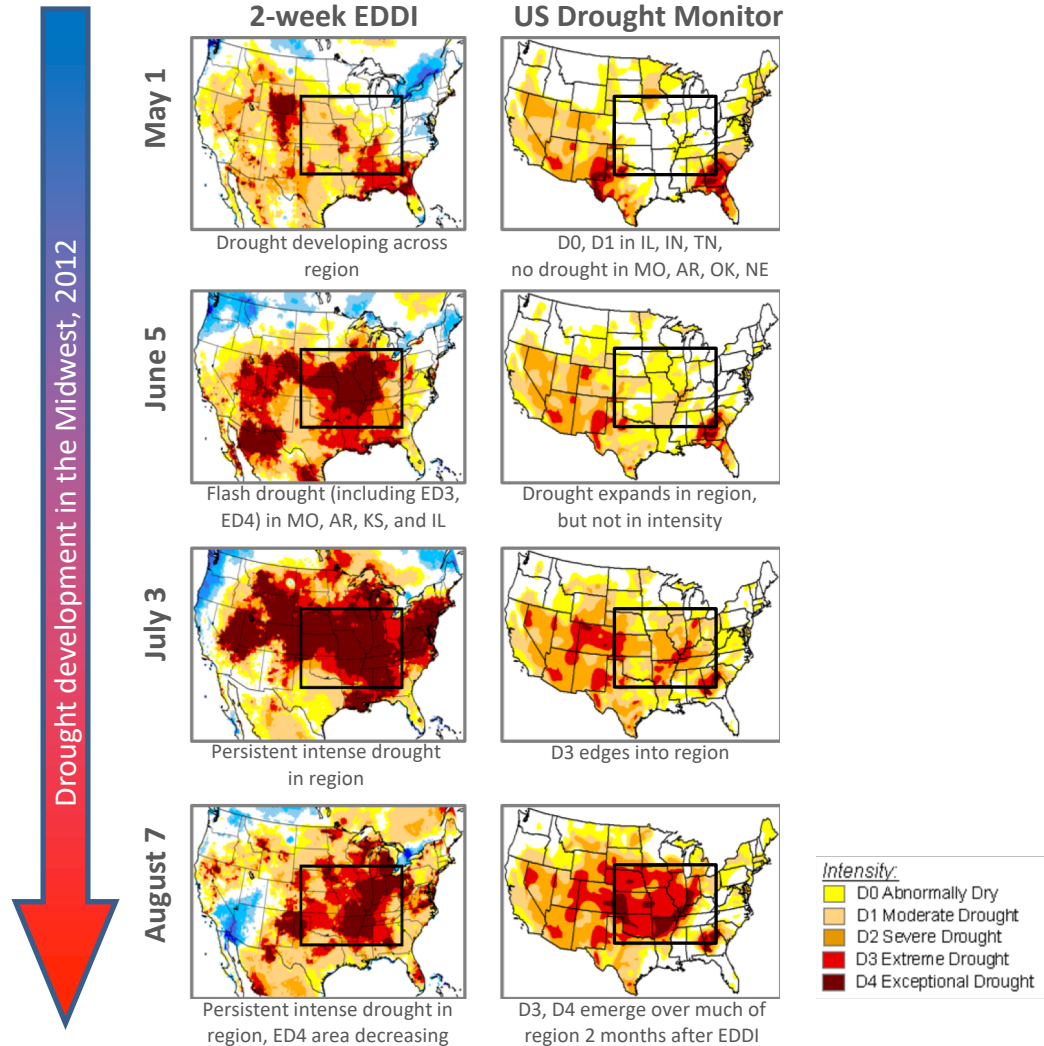


ECOLOGICAL
DROUGHT

EDDI – flash drought / early warning

Midwest US: 2012

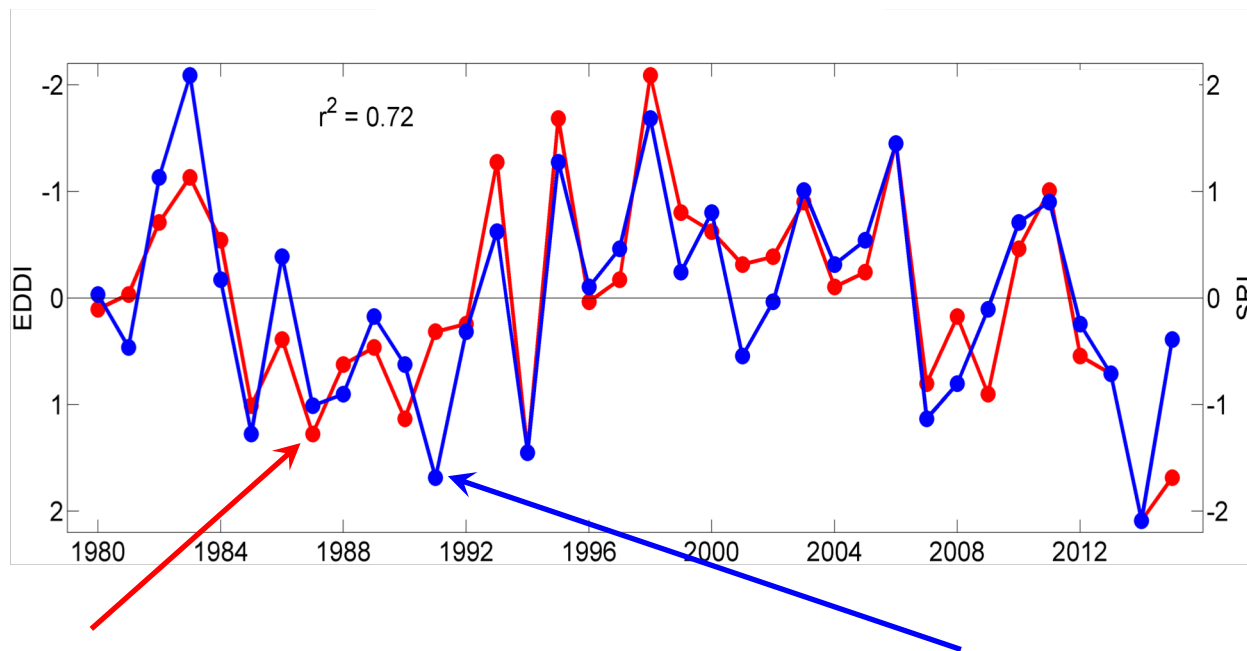
EDDI captured severe drought conditions ~2 months before the US Drought Monitor



EDDI – early warning of hydrological drought

*Can EDDI help predict
late-summer (low-flow)
streamflow?*

Sacramento River Basin, CA



6-month EDDI
(Nov-Apr)

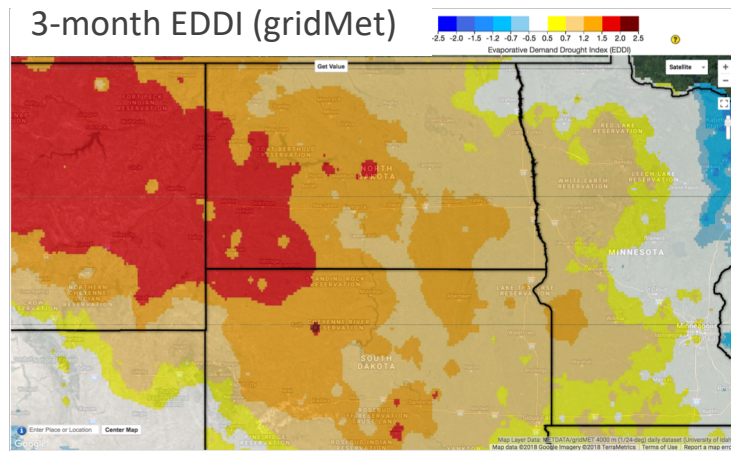
12-month SRI
(Oct-Sep)

SRI = Standardized Runoff Index

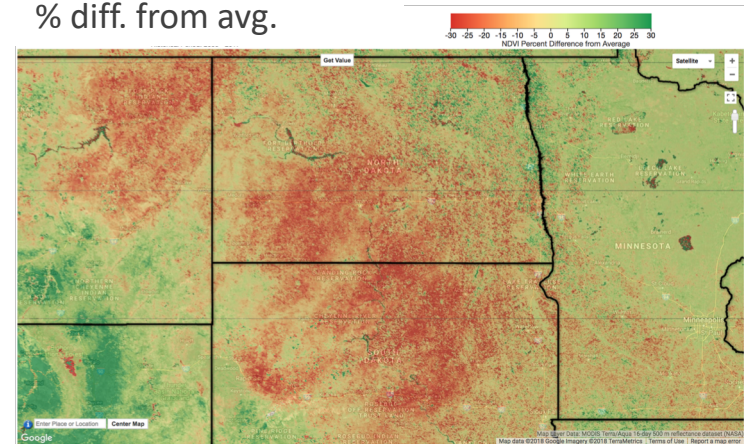
EDDI – complementing remote sensing

*Understanding remote sensing
anomalies of land surface
temperature, vegetation, and ET*

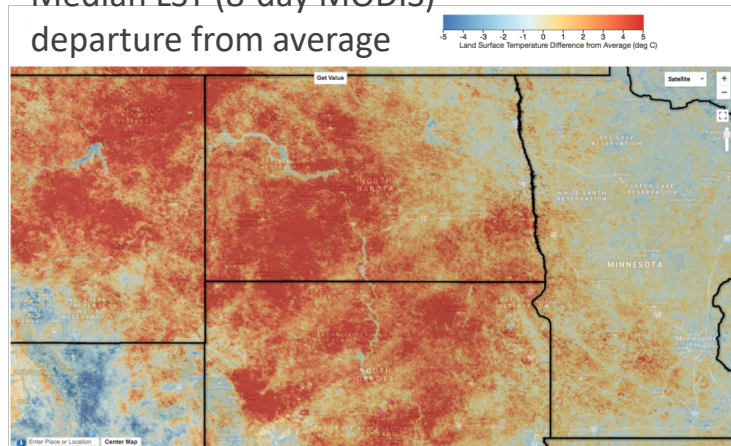
Northern Great Plains: May – July, 2017



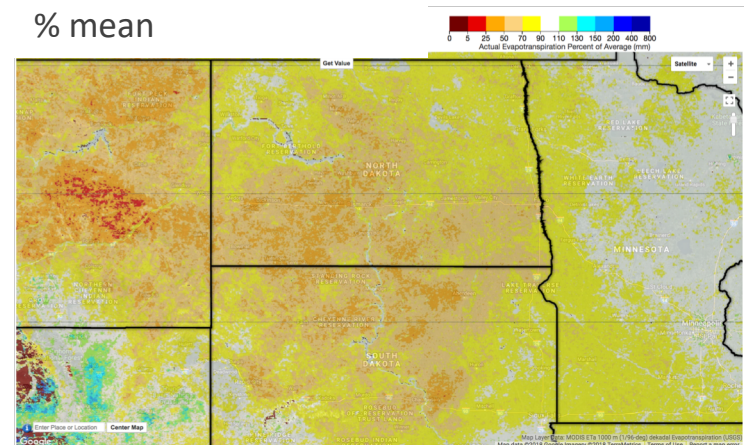
Median NDVI (16-day MODIS)
% diff. from avg.



Median LST (8-day MODIS)
departure from average



3-month ET (SSEBop)
% mean



Attribution – diagnosing drought's demand side

EDDI across 6-month period of drought

Northern Great Plains: 2017

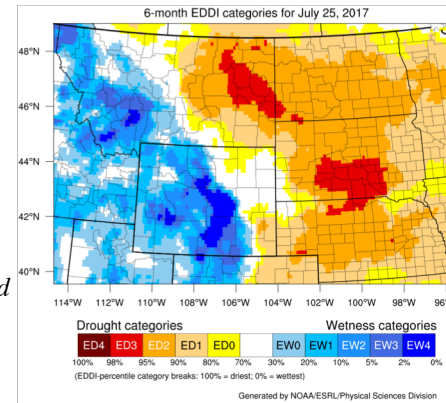
How much are
changes in E_0 due to
each driver's changes?

$$E_0 = f(T, R_d, q, U_2), \text{ so}$$

Decomposition of 6-month E_0 anomaly, as of July 25, 2017, $\Delta E_0 = \frac{\partial E_0}{\partial T} \Delta T + \frac{\partial E_0}{\partial R_d} \Delta R_d$

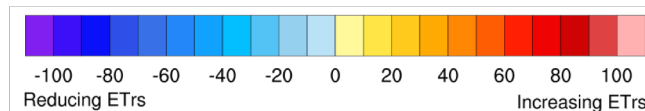
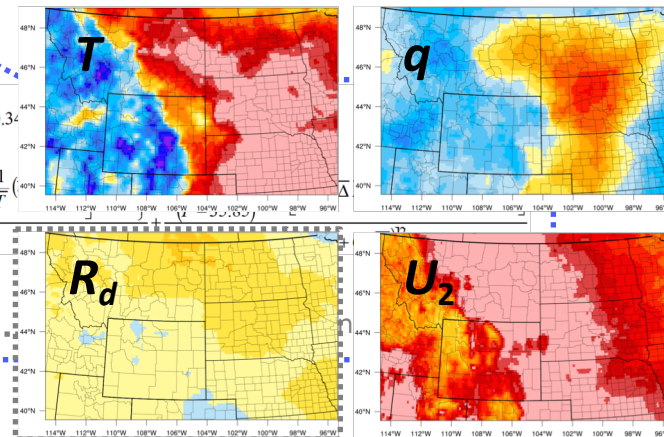
$$\Delta E_0 = \frac{\partial E_0}{\partial T} \Delta T + \frac{\partial E_0}{\partial R_d} \Delta R_d + \frac{\partial E_0}{\partial q} \Delta q + \frac{\partial E_0}{\partial U_2} \Delta U_2$$

derived analytically



observed in
reanalyses

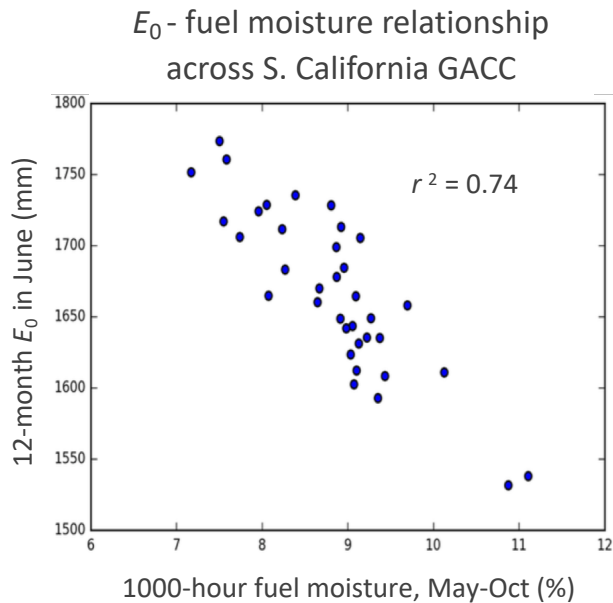
$$\frac{\partial E_{Tr}}{\partial T} = \left\{ \begin{aligned} &0.408 \Delta \left[\bar{R}_n \frac{4169.871 - 2\bar{T}}{(\bar{T} - 35.85)^2} - 4\sigma_{cd} (0.34 \right. \\ &\left. + \gamma C_n \frac{\bar{U}}{\bar{T}} \left[\frac{4098.171}{(\bar{T} - 35.85)^2} - \frac{1}{\bar{T}} \right] \right] \end{aligned} \right.$$



Contributions to 6-month E_0 anomaly (mm)

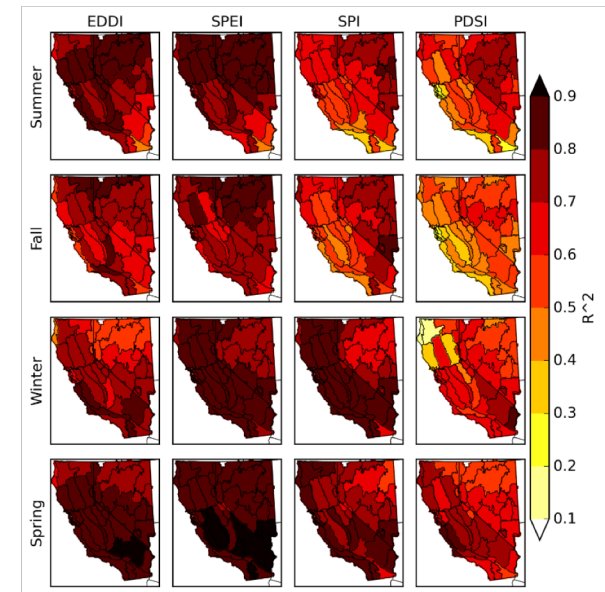
EDDI – wildfire risk monitoring

California and Nevada

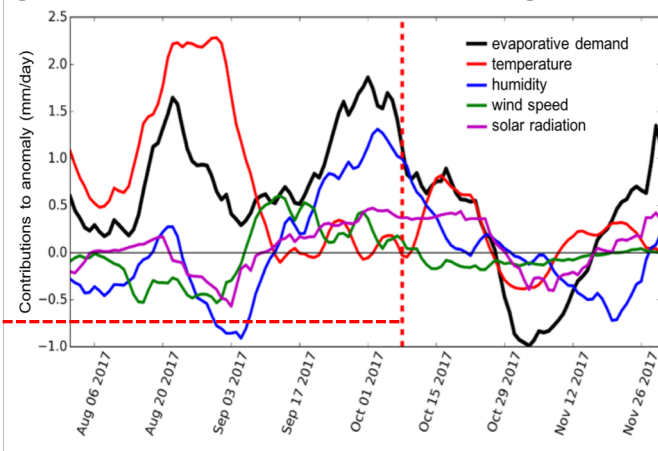


*Can EDDI provide early
warning of wildfire risk?*

Correlations to 1000-hour fuel moisture



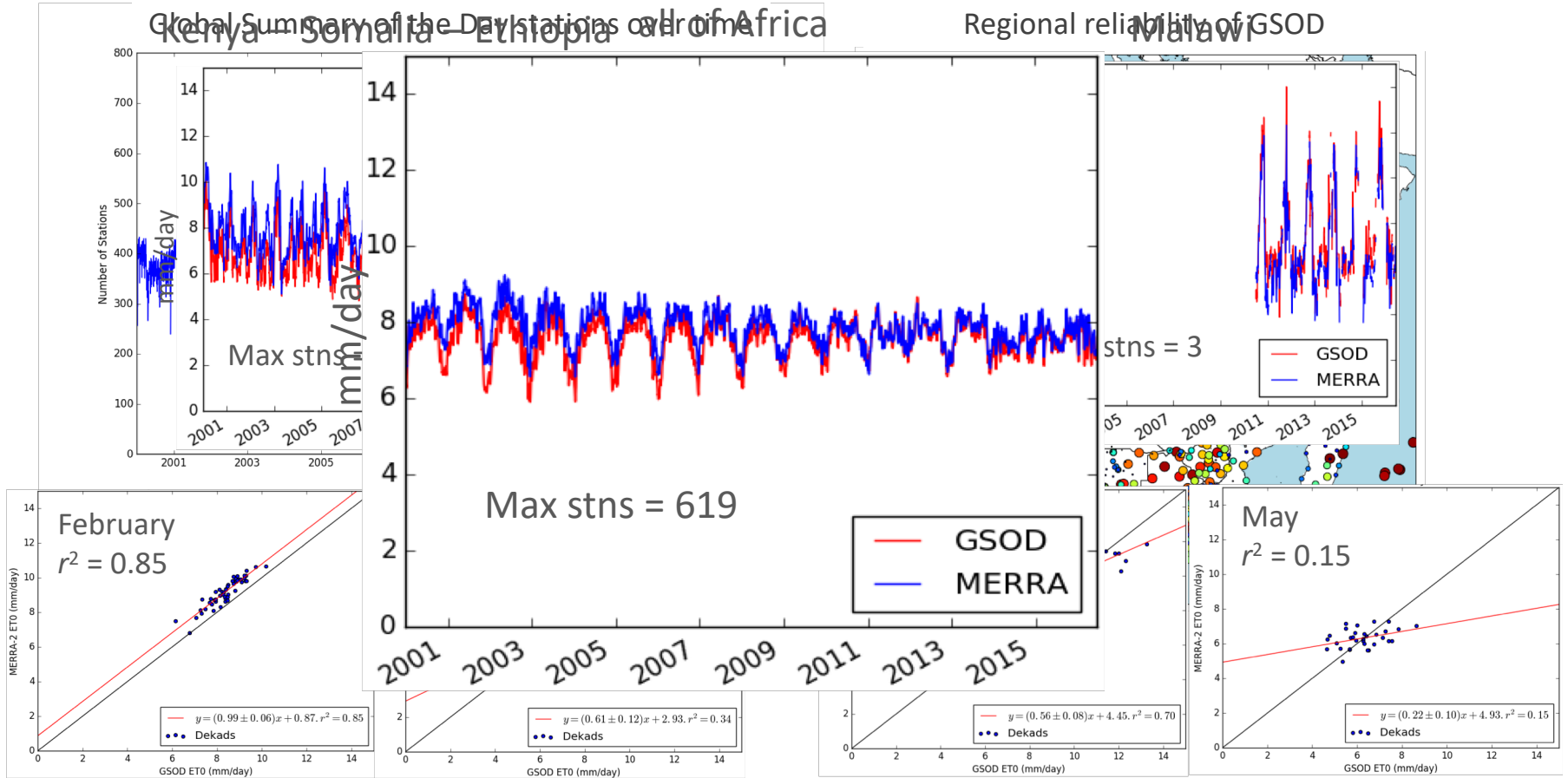
Diagnosis of ΔE_0 for Sonoma Co., CA Aug – Nov, 2017



Predicting fire-weather conditions for Tubbs fire:

- Decomposing EDDI into its drivers indicates a shift to a humidity-driven spike in E_0 about 4 weeks before the fire.

Verification of E_0 across Africa

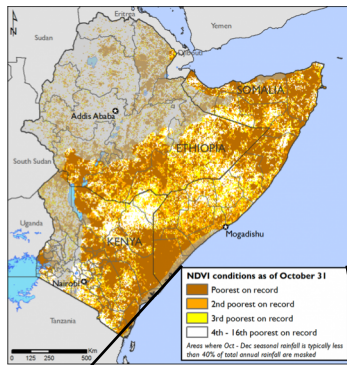


GSOD = Global Summary of the Day

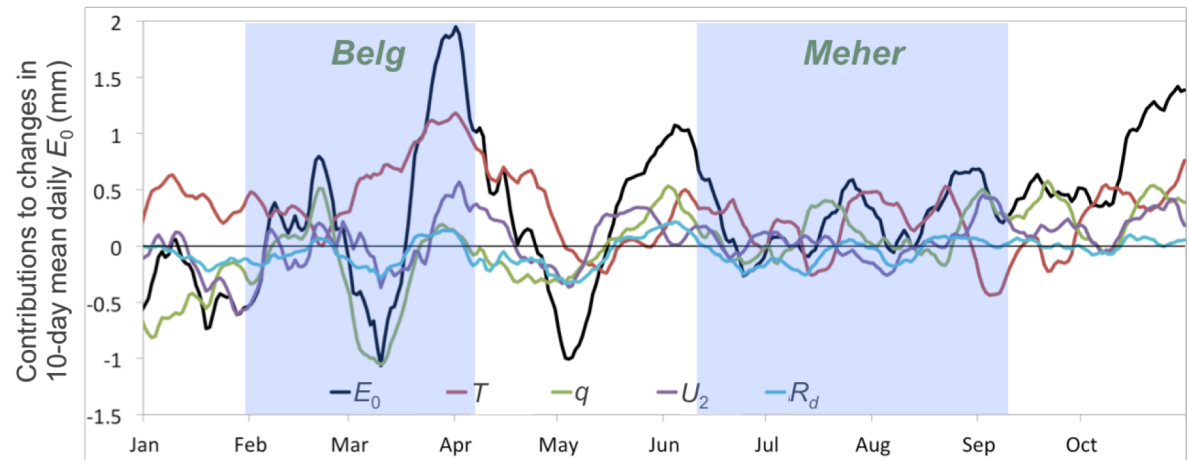
Attribution of demand side of drought

T = air temperature
 q = specific humidity
 U_2 = wind speed
 R_d = solar radiation

Horn of Africa, 2016 drought



Source: FEWSNET

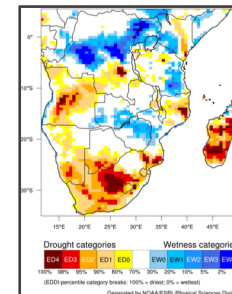
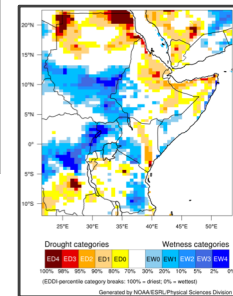
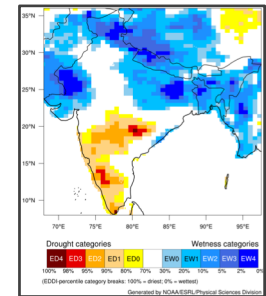
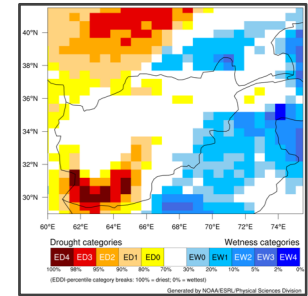
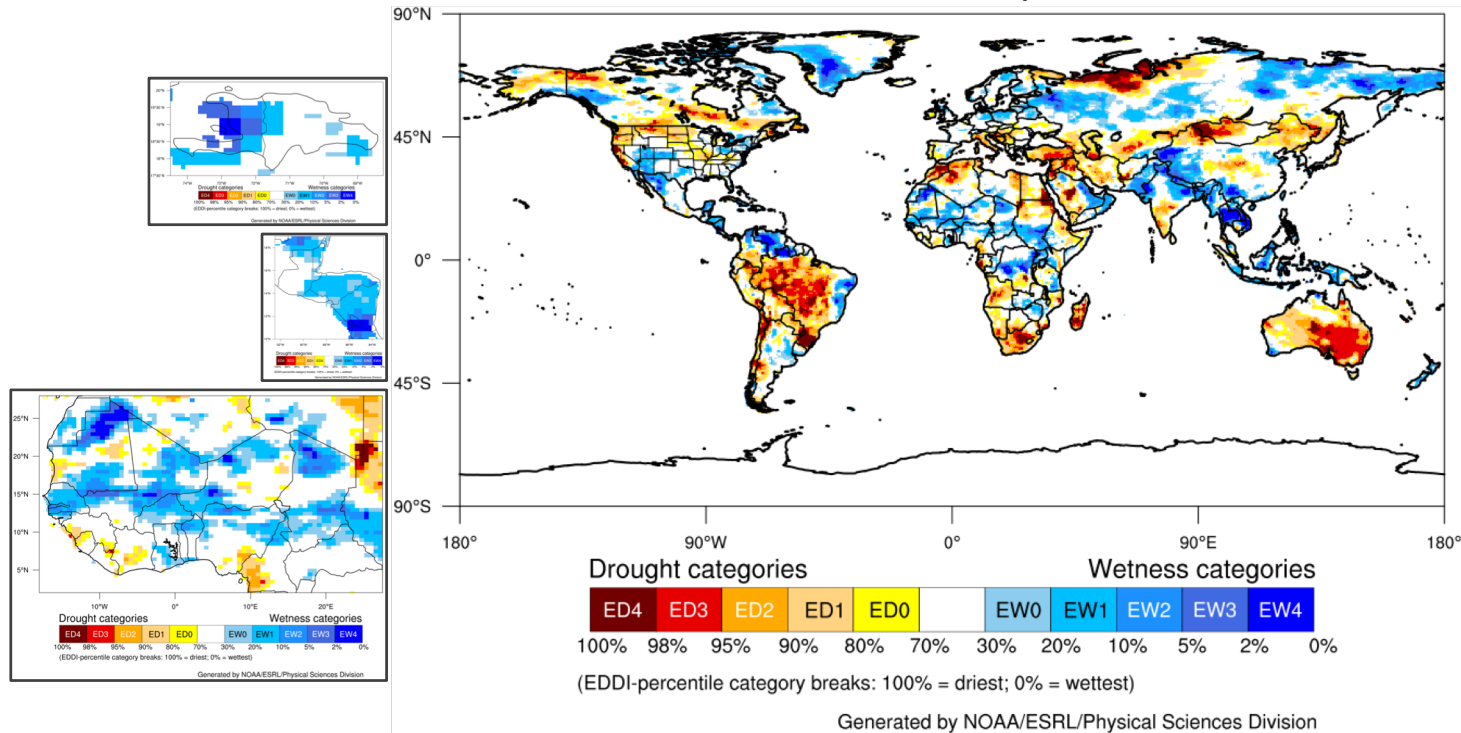


Changing E_0 dynamically forced by:

- elevated T ,
- Peaks coincide with elevated U_2 and/or q minima, troughs with q maxima.
- R_d plays little role

Global EDDI for famine early warning

1-month EDDI for July 31, 2017



Opportunities for collaboration

CONUS product:

- transition to National Water Model-forcing of EDDI (1 km, daily)
- examining drought linkages between NWM-derived soil moisture, evaporative demand, EDDI, and *ET*

Global product:

- data assimilation of reference *ET* observations
- improved downscaling of reference *ET* to 0.125°
- verification of EDDI in agricultural / hydrologic drought and early warning against independent R/S-based drought metrics (e.g., NDVI)